





It is shown that the NAO index is a good indicator of the strength of the NAO. The NAO index is defined as the difference in the sea level pressure (SLP) between the Azores and Iceland.

The NAO index is used to describe the variability of the NAO. The NAO index is defined as the difference in the sea level pressure (SLP) between the Azores and Iceland. The NAO index is used to describe the variability of the NAO. The NAO index is defined as the difference in the sea level pressure (SLP) between the Azores and Iceland.

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2. DATA AND METHODOLOGY

The data used in this study are the monthly mean sea level pressure (SLP) data from the National Centers for Environmental Prediction (NCEP) reanalysis project. The data cover the period from 1958 to 1996. The data are used to calculate the NAO index.

et al., 1996). It is shown

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Angelier and Ghil (1990), and Ghil and Angelier (1992) have shown that the NAO index is related to the North Atlantic temperature (NAT) index. The NAT index is defined as the difference in temperature between the Azores and Iceland.

The 500 Pa geopotential height (GHG) and geopotential height (GHGN) are defined as:

$$GHG = \frac{Z(\phi_0) - Z(\phi_s)}{\phi_0 - \phi_s} \tag{1}$$

$$GHGN = \frac{Z(\phi_n) - Z(\phi_0)}{\phi_n - \phi_0} \tag{2}$$

where  $\phi_n = 80^\circ N + \Delta$ ,  $\phi_0 = 60^\circ N + \Delta$ ,  $\phi_s = 40^\circ N + \Delta$  and  $\Delta = -2.5^\circ, 0^\circ, 2.5^\circ$ .

As shown in Figure 2, the 500 Pa geopotential height (GHG) and geopotential height (GHGN) are defined as:

Figure 4. A comparison of the 500 Pa geopotential height (GHG) and geopotential height (GHGN) for the NAO index. (a) GHG > 0, (b) GHG < -10 m.

GHG > 0

GHGN < -10 m

As shown in Figure 4, the 500 Pa geopotential height (GHG) and geopotential height (GHGN) are defined as:

et al. (1997) and

Figure 5 shows the 500 Pa geopotential height (GHG) and geopotential height (GHGN) for the NAO index.

Figure 6 shows the 500 Pa geopotential height (GHG) and geopotential height (GHGN) for the NAO index.

Figure 7 shows the 500 Pa geopotential height (GHG) and geopotential height (GHGN) for the NAO index.

Figure 8 shows the 500 Pa geopotential height (GHG) and geopotential height (GHGN) for the NAO index.

Figure 9 shows the 500 Pa geopotential height (GHG) and geopotential height (GHGN) for the NAO index.

Figure 10 shows the 500 Pa geopotential height (GHG) and geopotential height (GHGN) for the NAO index.

Figure 11 shows the 500 Pa geopotential height (GHG) and geopotential height (GHGN) for the NAO index.

Figure 12 shows the 500 Pa geopotential height (GHG) and geopotential height (GHGN) for the NAO index.

Figure 13 shows the 500 Pa geopotential height (GHG) and geopotential height (GHGN) for the NAO index.

Figure 14 shows the 500 Pa geopotential height (GHG) and geopotential height (GHGN) for the NAO index.

et al., 1998).

Table 1. The NAO index and the North Atlantic temperature (NAT) index for the period 1962-1996.

NAO	1962	1963	1964	1965	1966	1969	1970	1979	1996
NAT	1973	1978	1983	1989	1990	1992	1993	1994	1995

The NAO index is defined as the difference in SLP between the Azores and Iceland. The NAT index is defined as the difference in temperature between the Azores and Iceland.

3. THE STATISTICAL RELATIONSHIP

A positive NAO is associated with a strong westerly flow of air from the Azores to the British Isles, resulting in a mild and wet winter. In contrast, a negative NAO is associated with a weak westerly flow, resulting in a cold and dry winter.

*et al.*, 1997). Using NCEP reanalysis data, we have examined the relationship between the NAO and blocking patterns over the North Atlantic region.

Figure 1 shows the time series of the NAO index and the number of blocking events over the North Atlantic region from 1960 to 1990. The NAO index is defined as the difference in sea level pressure between the Azores and the British Isles. The number of blocking events is defined as the number of days when a blocking pattern is present over the North Atlantic region.

Figure 1 shows that the NAO index and the number of blocking events are highly correlated. The correlation coefficient is 0.82. This indicates that a strong NAO is associated with a high number of blocking events, while a weak NAO is associated with a low number of blocking events.

Figure 2 shows the time series of the NAO index and the number of blocking events over the North Atlantic region from 1958 to 1996. The NAO index is defined as the difference in sea level pressure between the Azores and the British Isles. The number of blocking events is defined as the number of days when a blocking pattern is present over the North Atlantic region.

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- Cd 130°W
- Ah 60°WE;
- Ep 10°80E.

Figure 2 shows that the NAO index and the number of blocking events are highly correlated. The correlation coefficient is 0.82. This indicates that a strong NAO is associated with a high number of blocking events, while a weak NAO is associated with a low number of blocking events.

$$\times 4 = 152 \text{ (38 grid boxes). The correlation coefficient is } -0.45 \text{ (Figure 1)}$$

Figure 3 shows the time series of the NAO index and the number of blocking events over the North Atlantic region from 1958 to 1996. The NAO index is defined as the difference in sea level pressure between the Azores and the British Isles. The number of blocking events is defined as the number of days when a blocking pattern is present over the North Atlantic region.



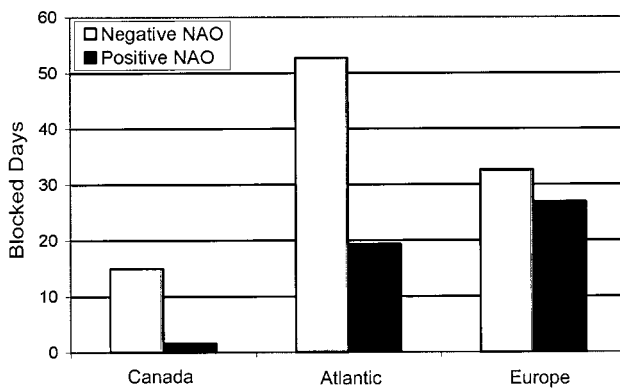


Fig. 3. Comparison of NAO (a) NAO (b) Td  
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Fig. 4. (A) NAO index and (B) North Atlantic Index  
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 The NAO index in Fig  
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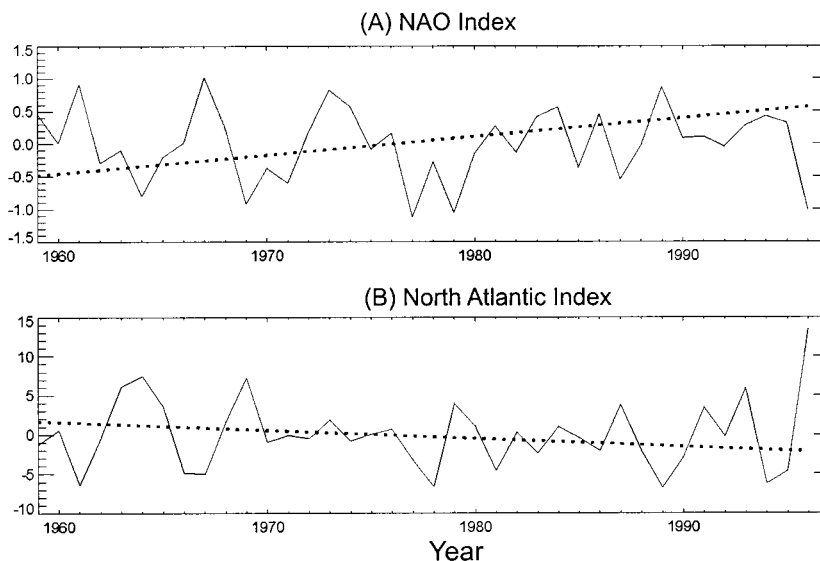


Fig. 4. (A) NAO index and (B) North Atlantic Index

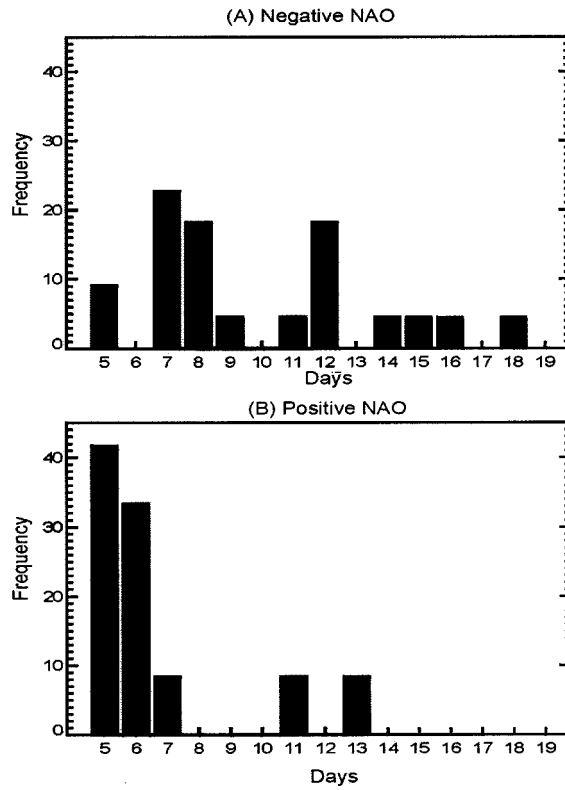


Fig. 5. Frequency of days for

In the NAO, 15% of the days are classified as negative NAO (6 of Dec (1989) and 10 of Oct (Fig. 5) and 18 of the days are classified as positive NAO (42 of Dec (1989) and 33 of Oct (Fig. 5)).

The 500 Pa NAO index is used to identify the NAO. The NAO index is defined as the difference between the 500 Pa geopotential height at 0° and 60°W. The NAO index is used to identify the NAO. The NAO index is used to identify the NAO. The NAO index is used to identify the NAO.

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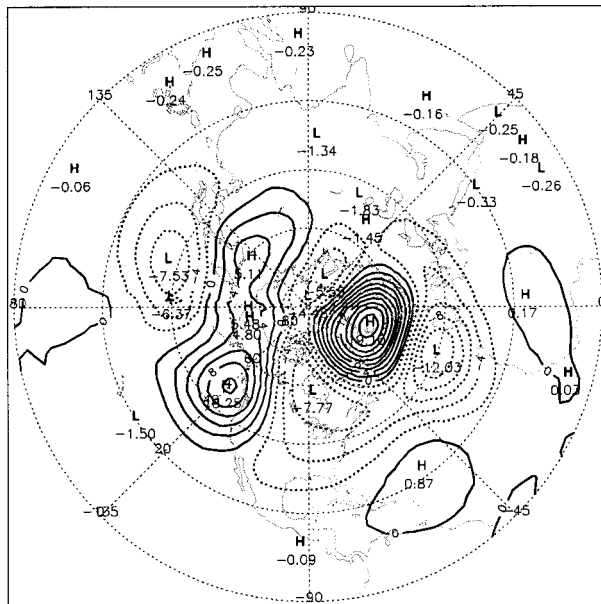


Fig. 6. Blocking Index at 500 hPa for the NAO in the period 1950-2000. Contour interval is 0.25.

#### 4. THE IMPACT OF THE NAO ON BLOCKING

The NAO is a dominant mode of climate variability in the North Atlantic region. It is characterized by a seesaw in atmospheric pressure between the Azores and Iceland. The NAO index (NAOI) is defined as the difference in sea level pressure (SLP) between the Azores and Iceland. The NAOI is a measure of the strength of the NAO. The NAOI is a measure of the strength of the NAO. The NAOI is a measure of the strength of the NAO.

Following Davis (1979), the NAOI is defined as the difference in SLP between the Azores and Iceland.

$$\frac{\partial}{\partial t} \left( \nabla^2 \psi - \frac{\psi}{\lambda^2} \right) + J(\psi, \nabla^2 \psi + h) + \beta \frac{\partial \psi}{\partial x} = k \nabla^2 (\psi^* - \psi) \quad (3)$$

where  $\psi$  is the stream function,  $h$  is the height,  $J$  is the Jacobian,  $\beta = L/a \cos \phi_0$ ,  $\phi_0$  is the latitude,  $a$  is the radius of the Earth,  $k = D_E/2H$ ,  $D_E$  is the eddy diffusion coefficient,  $H$  is the scale height,  $\psi^*$  is the stream function of the NAO, and  $\lambda$  is the wavelength of the NAO. The NAOI is a measure of the strength of the NAO.

$$F_A = \sqrt{2} \sin y, \quad F_K = 2 \sin nx \sin y, \quad F_L = 2 \sin nx \sin y$$

Substituting Eq. (4) into Eq. (1) and using the boundary conditions (2), we obtain

$$\psi = \psi_A F_A + \psi_K F_K + \psi_L F_L$$

$$\psi^* = \psi_A^* F_A + \psi_K^* F_K$$

$$h = \frac{h_0}{2H} F_K$$

where

$$\dot{\psi}_A = -k(\psi_A - \psi_A^*) + h_{01}\psi_L \tag{4}$$

$$\dot{\psi}_K = -(\alpha_{n1}\psi_A - \beta_{n1})\psi_L - k(\psi_K - \psi_K^*) \tag{5}$$

$$\dot{\psi}_L = -(\alpha_{n1}\psi_A - \beta_{n1})\psi_K - k(\psi_K - h_{n1})\psi_A \tag{6}$$

where

$$\gamma_{n1} = \frac{8\sqrt{2}}{3\pi} n, \quad h_{01} = \gamma_{n1} \frac{h_0}{2H}, \quad h_{n1} = \frac{\gamma_{n1}}{n^2 + 1} \frac{h_0}{2H}, \quad \alpha_{n1} = \frac{n^2}{n^2 + 1} \gamma_{n1} \quad \text{and} \quad \beta_{n1} = \frac{n}{n^2 + 1} \beta$$

Substituting Eq. (4) into Eq. (6), we obtain

$$\psi_A^3 + v_1 \psi_A^2 + v_2 \psi_A + v_3 = 0 \tag{7}$$

where

$$v_1 = -(2\beta_{n1}\alpha_{n1} + \alpha_{n1}^2 \psi_A^*) / \alpha_{n1}^2,$$

$$v_2 = (\beta_{n1}^2 + k^2 + 2\beta_{n1}\alpha_{n1}\psi_A^* + h_{01}(h_{n1} - \alpha_{n1}\psi_K^*)) / \alpha_{n1}^2$$

and

$$v_3 = ((\beta_{n1}^2 + k^2)\psi_A^* + h_{01}\beta_{n1}\psi_K^*) / \alpha_{n1}^2$$

The root of Eq. (7) is given by the Cardan formula (Chandrasekhar 1979).

The bifurcation diagram

is shown in Fig. 7. The parameters are  $k = 2 \times 10^{-2}$ ,  $h_0/H = 0.05$ ,  $n = 2$ ,  $\psi_A^* = 0.2$ .

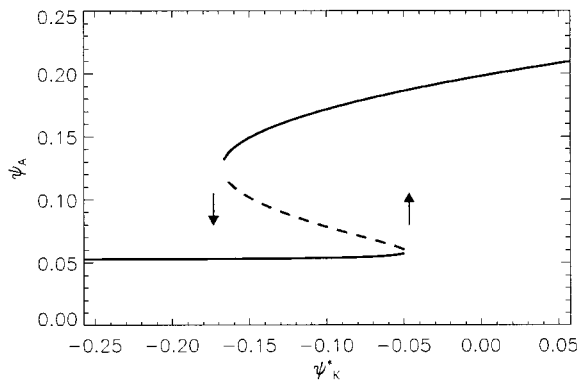


Fig. 7. Bifurcation diagram of  $\psi_A$  versus  $\psi_K^*$ .

The parameters are  $k = 2 \times 10^{-2}$ ,  $h_0/H = 0.05$ ,  $n = 2$ ,  $\psi_A^* = 0.2$ .

$\psi_K^*$ .

is characterized by a stable low index equilibrium (Fig. 7). In this case, the circulation is characterized by a strong westerly flow in the North Atlantic, with a deep low pressure cell over the North Atlantic and a deep high pressure cell over the Azores. The circulation is characterized by a strong westerly flow in the North Atlantic, with a deep low pressure cell over the North Atlantic and a deep high pressure cell over the Azores.

$$\psi_k^* > -0.05), \text{ which is}$$

characterized by a stable high index equilibrium (Fig. 8). In this case, the circulation is characterized by a strong westerly flow in the North Atlantic, with a deep low pressure cell over the North Atlantic and a deep high pressure cell over the Azores. The circulation is characterized by a strong westerly flow in the North Atlantic, with a deep low pressure cell over the North Atlantic and a deep high pressure cell over the Azores.

$$-0.17 < \psi_A^* < -0.05), \text{ and}$$

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$$\psi_k^* < -0.17), \text{ which is}$$

characterized by a stable high index equilibrium (Fig. 8). In this case, the circulation is characterized by a strong westerly flow in the North Atlantic, with a deep low pressure cell over the North Atlantic and a deep high pressure cell over the Azores. The circulation is characterized by a strong westerly flow in the North Atlantic, with a deep low pressure cell over the North Atlantic and a deep high pressure cell over the Azores.

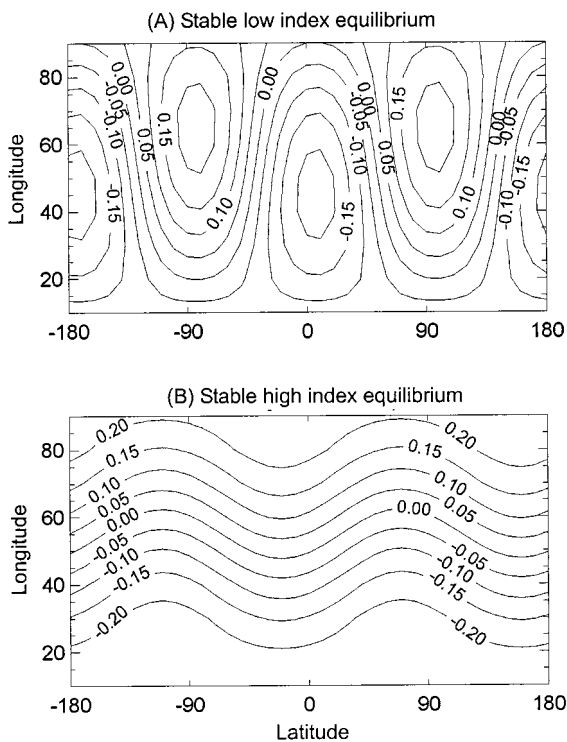
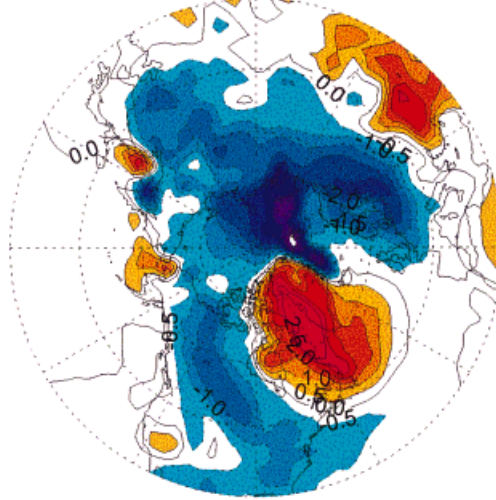


Fig. 8. Same as Fig. 7.

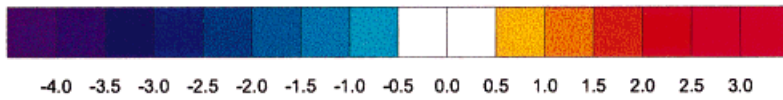
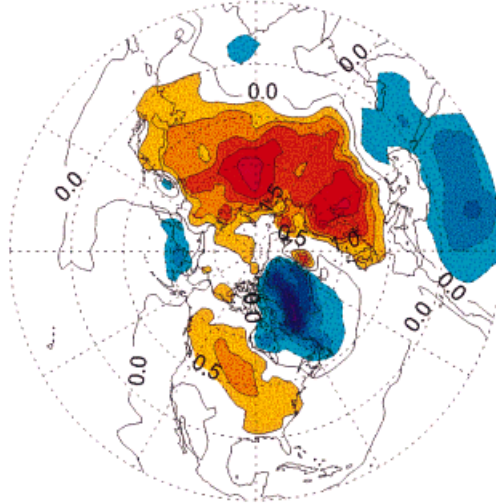
$$\psi_k^* = -0.1 \text{ in Fig. 7, which is}$$



(A) Negative NAO



(B) Positive NAO



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5. DISCUSSION AND CONCLUSIONS

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Conceptual Model for Blocking

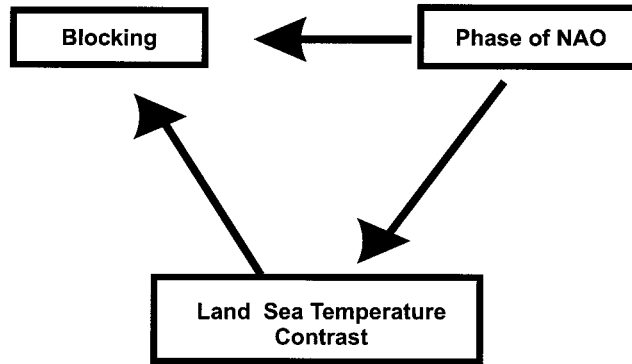


Fig. 9. A conceptual model for blocking.

Blocking is a phenomenon that occurs in the North Atlantic region. It is characterized by a high-pressure system over the North Atlantic and a low-pressure system over the Azores. This configuration leads to a northward shift of the storm track, resulting in a wet and mild winter in Europe. The North Atlantic Oscillation (NAO) is a climate mode that describes the variability in the pressure difference between the Azores and Iceland. The NAO has two main phases: the positive phase and the negative phase. In the positive phase, the pressure difference is large, leading to a strong and stable jet stream. In the negative phase, the pressure difference is small, leading to a weak and unstable jet stream. The NAO is influenced by various factors, including the Land-Sea Temperature Contrast (LSTC). The LSTC is the difference in temperature between the land and the sea. In the North Atlantic, the LSTC is large due to the presence of the Gulf Stream, which warms the sea surface. This large LSTC leads to a strong and stable jet stream, which is characteristic of the positive phase of the NAO.

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